

REPLY

10.1002/2014JC010222

This article is a reply to  
 McDougall and Barker [2014],  
 doi:10.1002/2014JC010066.

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Citation:

Gerkema, T., H. L. Swinney, M. Stone,  
 and M. Marder (2014), Reply to  
 comment by McDougall and Barker on  
 "Buoyancy frequency profiles and  
 internal semidiurnal tide turning  
 depths in the oceans," *J. Geophys. Res.*  
*Oceans*, 119, 9033–9034, doi:10.1002/  
 2014JC010222.

Received 9 JUN 2014

Accepted 19 NOV 2014

Accepted article online 24 NOV 2014

Published online 30 DEC 2014

## Reply to comment by McDougall and Barker on "Buoyancy frequency profiles and internal semidiurnal tide turning depths in the oceans"

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### 1. Introduction

In a recent paper [King *et al.*, 2012], we studied the turning depths of internal tides in the lower layers of the ocean. The turning depth is determined by latitude, wave frequency, and the buoyancy frequency  $N$ . Exploring all available WOCE data, we adopted King *et al.* [2012, equations (2) and (3)] for the calculation of  $N^2$ . These are the same as equations (2a) and (2c) in the Comment by McDougall and Barker [2014], which they present as correct expressions for  $N^2$ . Thus McDougall and Barker [2014] challenges none of the results presented in King *et al.* [2012], which were obtained using scripts from TEOS-10, as listed in Table 1 of our paper.

In King *et al.* [2012, section 2.2], we also made a critical remark about the use of potential density in determining  $N^2$ , but we mistakenly grouped the CSIRO-script `sw_bfrq.m` with the inexact ones. In the following section, we clarify the intent of King *et al.* [2012, section 2.2].

Our misinterpretation of the script `sw_bfrq.m` arose from its giving in the preamble Eq. 3.7.15 for  $N^2$  from Gill [1982, p. 54], which adopts a constant reference level for potential temperature. McDougall and Barker [2014] clarifies that the method actually used in `sw_bfrq.m` is a different one, namely the adiabatic leveling method, which correctly uses local reference levels, and that the citation in the preamble of `sw_bfrq.m` should have been to Millard *et al.* [1990]. We thank McDougall and Barker for this clarification, and we apologize for our erroneous statement about the correctness of `sw_bfrq.m`. Further, we regret not contacting the authors of the script prior to the publication of King *et al.* [2012].

### 2. Potential Density and the Adiabatic Leveling Method

In general, potential density may be defined with respect to any reference pressure,  $p_r$ . For given temperature  $T$ , (absolute) salinity  $S$  and pressure  $p$ , we then write the density a water parcel would have if brought adiabatically from its initial level  $p$  to reference pressure  $p_r$ , i.e., its potential density, as  $\rho_r(p, T, S, p_r)$ . For given static vertical profiles  $T(p)$  and  $S(p)$ , we introduce the short-hand notation  $\rho_r(p, p_r)$ , for convenience.

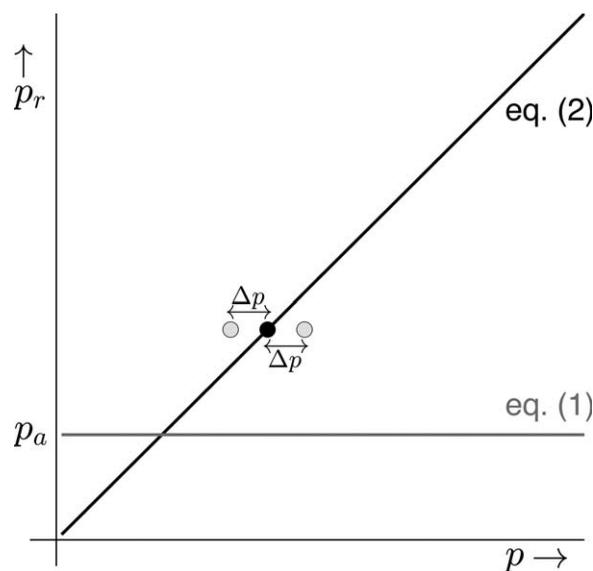
There are now two fundamentally different cases to be distinguished. If the partial derivative is taken with respect to the vertical coordinate (for which hydrostatic pressure  $p$  serves as a proxy) and the reference pressure is regarded as a constant, say atmospheric value  $p_r = p_a$ , the resulting expression is not proportional to  $N^2$ :

$$g^2 \frac{\partial \rho_r}{\partial p}(p, p_a) \neq N^2. \tag{1}$$

The sole aim of King *et al.* [2012, section 2.2] was to reiterate this point, which was made before by, e.g., Gill [1982, p. 55].

The other case (not discussed in King *et al.* [2012]) is when  $p_r$  is regarded as a variable, and is afterward assigned the value  $p_r = p$ . The resulting expression then is an exact one for local gravitational stability:

$$N^2 = g^2 \left( \frac{\partial \rho_r}{\partial p}(p, p_r) \right) \Big|_{p_r=p}. \tag{2}$$



**Figure 1.** Plane spanned by the independent variables on which the generalized potential density  $\rho_r$  depends: initial pressure  $p$  and reference pressure  $p_r$ . Usually, the latter is taken constant (e.g., at atmospheric pressure), giving the horizontal line at  $p_r = p_a$  (in grey), adopted in the evaluation of the left-hand side of (1). If the partial derivative with respect to  $p$  is evaluated along the diagonal  $p_r = p$  (in black), the result is the identity (2). The circles sketch the discretized way of evaluation proposed by Millard *et al.* [1990].

It is crucial here to take the partial derivative *before* assigning the value to  $p_i$ ; after all, changing the order would simply result in the vertical gradient of in situ density  $\rho$ , since  $\rho_r(p, p) = \rho(p)$ . Although written differently, equation (2) is equivalent to equation (10) in McDougall and Barker [2014].

The difference between how (1) and (2) are evaluated can be illustrated as follows. We regard the independent variables  $p$  and  $p_r$  in plane coordinates (Figure 1). For potential density with a constant reference level (on the ordinate,  $p_a$ , say), one follows the partial derivative with respect to  $p$  along a horizontal line. This is the gradient of potential density in the usual sense (i.e., with constant reference level), which is *not* an exact

measure of static stability, see (1). If, on the other hand, one follows the partial derivative with respect to  $p$  along the *diagonal*  $p_r = p$ , then one finds the correspondence to (2). This is not a ‘gradient’ in the common sense of the word, but it does represent an exact expression for  $N^2$ . The potential density is now a *local* one, involving a continuum of reference levels.

The “adiabatic leveling method” proposed by Millard *et al.* [1990] casts (2) into a discretized form. At any point on the diagonal (black circle in Figure 1), one evaluates the horizontal partial derivative from a central-difference method using neighboring points at both sides on the horizontal (open circles), so that the right-hand side of (2) is approximated by

$$g^2 \frac{\rho_r(p + \Delta p, p) - \rho_r(p - \Delta p, p)}{2\Delta p}.$$

### 3. Conclusion

McDougall and Barker [2014] argues for the validity of the adiabatic leveling method for calculating the buoyancy frequency. However, this is not a point of contention; the adiabatic leveling method was not criticized by King *et al.* [2012], it was not even mentioned. As the previous section shows, we fully agree with the validity of that method. Our error was that we failed to realize that the script `sw_bfrq.m` uses that method. Again, we apologize for our misinterpretation of the script.

In this Reply we have not mentioned section 3 of McDougall and Barker [2014], since we fully agree with it. We should however point out that it concerns an older version of the equation of state, EOS-80, one that we did not use in King *et al.* [2012]. We used the modern standard TEOS-10 throughout.

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### Acknowledgments

The research at the University of Texas at Austin was supported by the Office of Naval Research MURI grant N000141110701. The authors acknowledge the AGU Publications Data Policy; this Reply to the Comment contains no data.